

# The International Height Reference System (IHRIS) and its realisation, the International Height Reference Frame (IHRF)

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# Outline



- Motivation
- Definition of the International Height Reference System (IHR)
- Realisation of the IHR: the International Height Reference Frame (IHRF)
  - Station selection for the reference network
  - Some considerations for the determination of IHR/IHRF coordinates
- Colorado experiment: comparison of potential values and learnings from a successful international cooperation initiative
- Participation of Latin America in the implementation of the IHR/IHRF

# Motivation

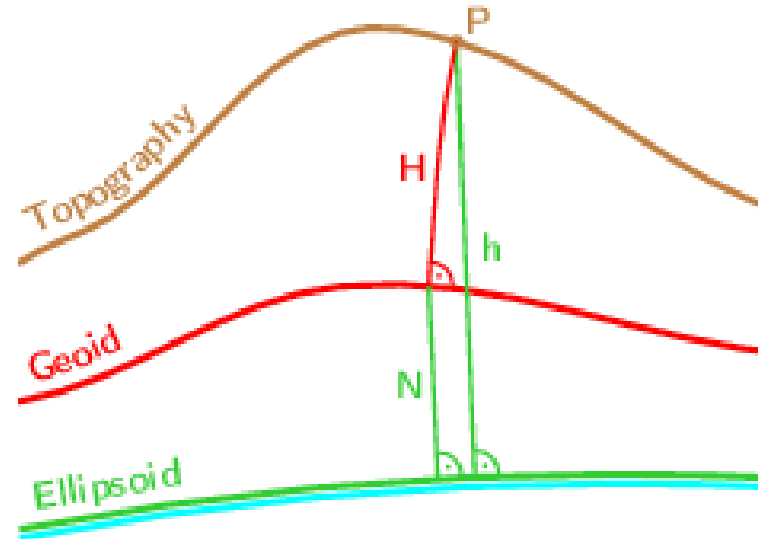
## 1) Vertical coordinates used in practice:

- $h$  → ellipsoidal heights (GNSS positioning);
- $H$  → Physical heights (levelling + gravity reductions);
- $N$  → (Quasi-)geoid undulations (gravity field modelling).

## 2) Everyone using GNSS positioning and requiring physical heights demands

$$H = h - N$$

with consistency at the cm-level and worldwide.



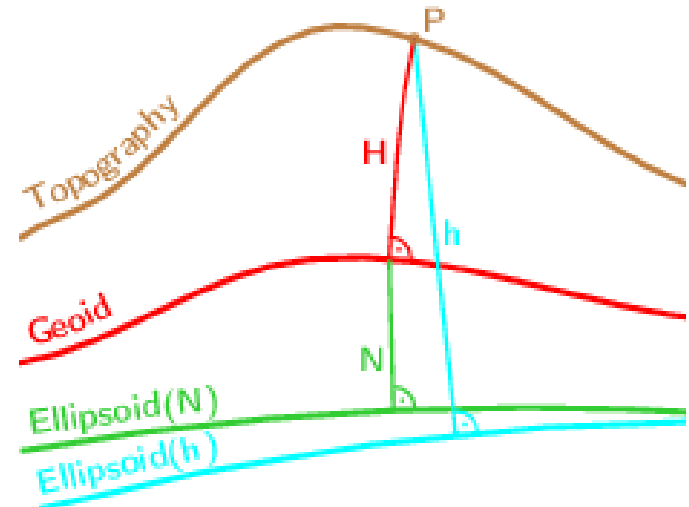
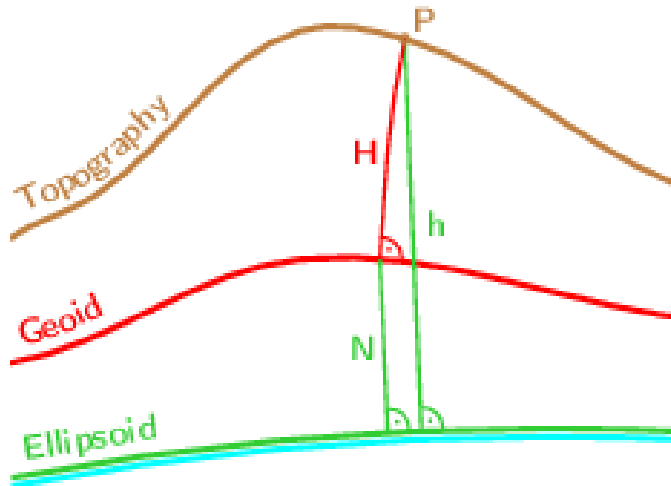
$H = h - N$  in theory ...

but in practice, e.g.



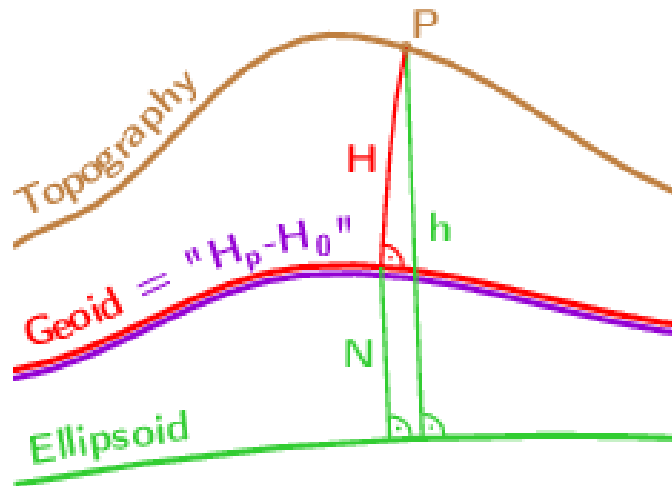
- Ellipsoidal heights  $h$  and geoid undulations  $N$  must be given w.r.t. the **same ellipsoid**:
  - $[X, Y, Z] \Leftrightarrow [\varphi, \lambda, h]$
  - Reference field (surface) for the geoid computation and for scaling global gravity models (GGM)

- **Different ellipsoid parameters** ( $a$ ,  $GM$ ) are used in geometry and gravity, for instance:
  - Geometric coordinates  $[\varphi, \lambda, h]$  referring to the GRS80 ellipsoid or to the WGS84 ellipsoid are practically identical
  - Geoid undulations  $N$  referring to the WGS84 ellipsoid present a discrepancy of about 93 cm w.r.t. geoid undulations referring to the GRS80



$H = h - N$  in theory ...

- Physical heights  $H$  and (quasi)geoid undulations  $N$  must reflect the same **reference surface**:
  - $H_p$  (from levelling) –  $H_0$  (datum point)
    - geoid from geometry
  - $N$  (from the GBVP)
    - geoid from gravity

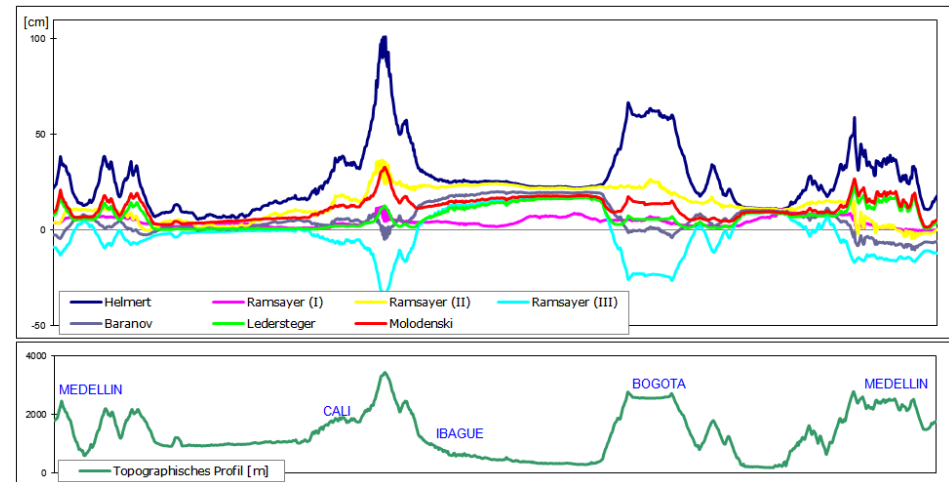


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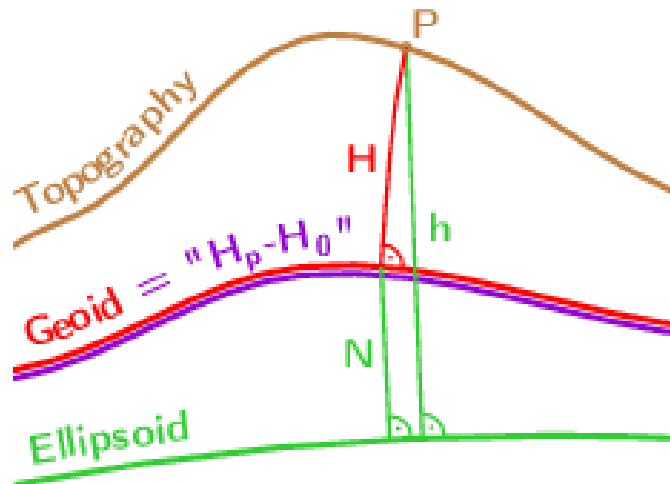
- Orthometric heights and gravimetric geoid use different hypotheses
- Different tide systems for  $H$  and  $N$
- Systematic errors over long distances in levelling (reliability of  $H_p - H_0$ )

Levelling-based physical heights with different gravity corrections



$H = h - N$  in theory ...

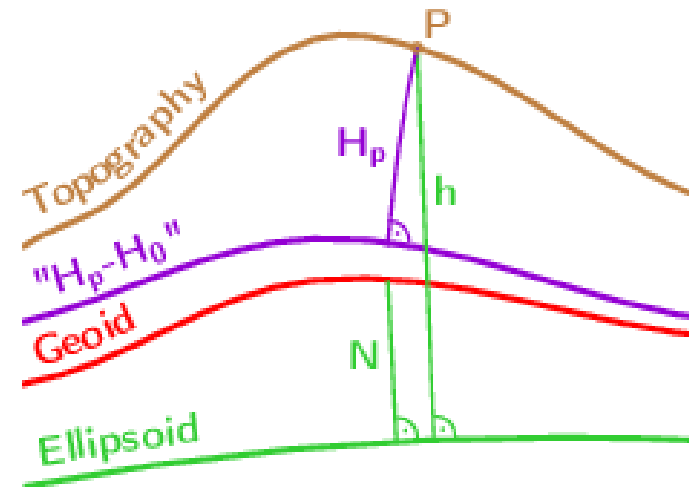
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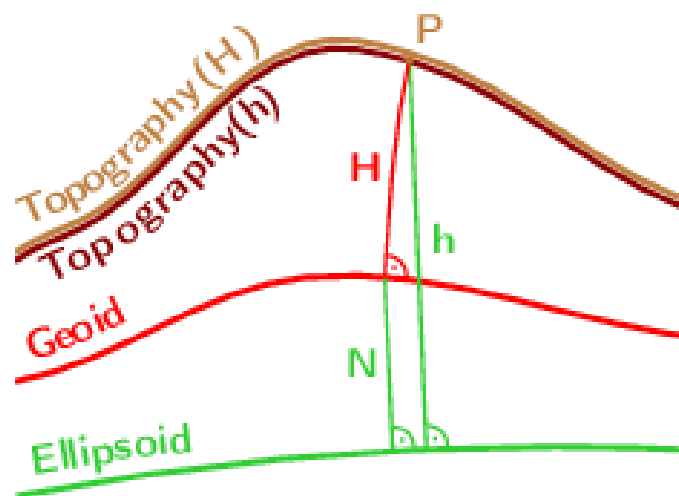
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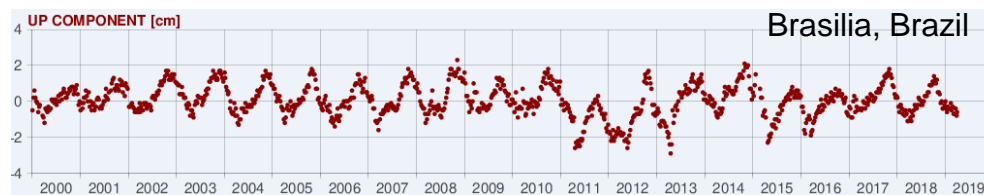
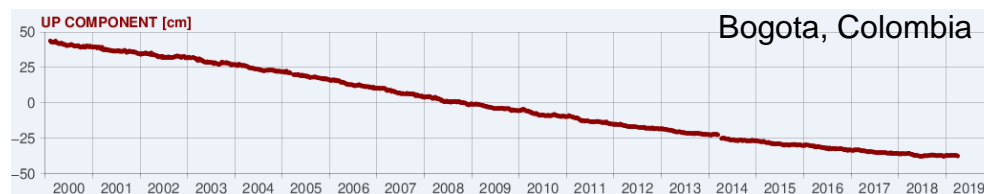


- Physical heights  $H$  and ellipsoidal heights  $h$  must represent the same Earth's surface

- Different reference epochs (with unknown  $dH/dt$ )
- Different reductions (Earth-, ocean-, atmospheric tides, ocean and atmospheric loading, post-glacial rebound, etc.)



Time series of ellipsoidal heights,



but levelling-based physical heights constant ( $dH/dt = 0$ )

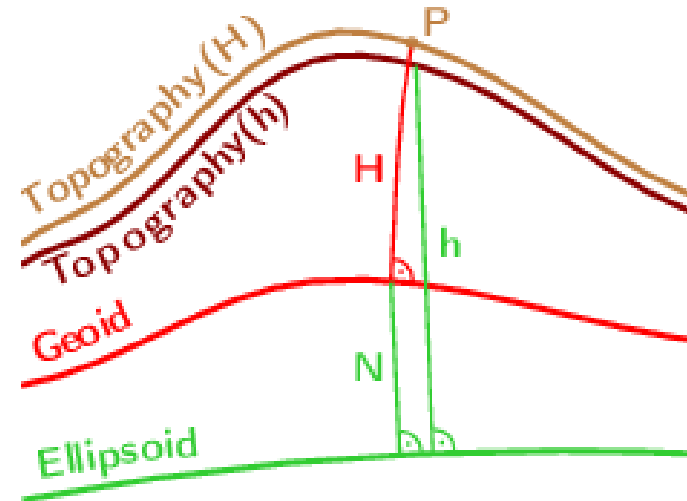
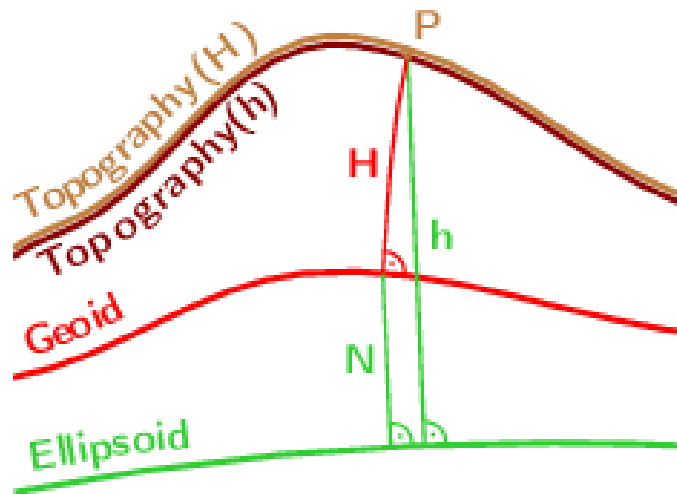
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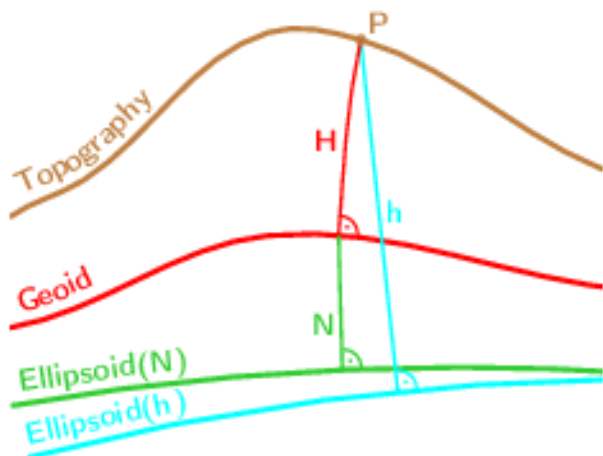
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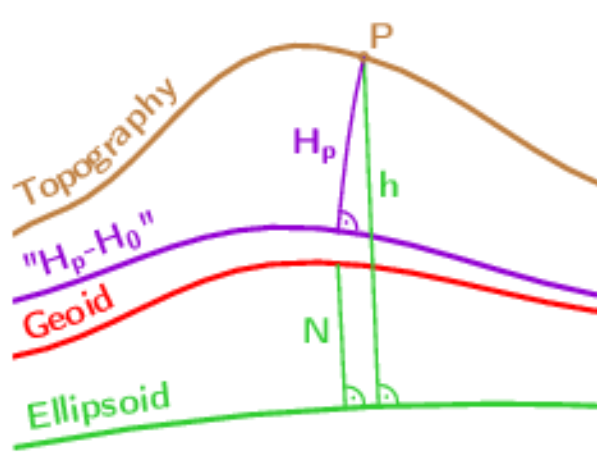


# A global unified height system is needed to ensure consistency between

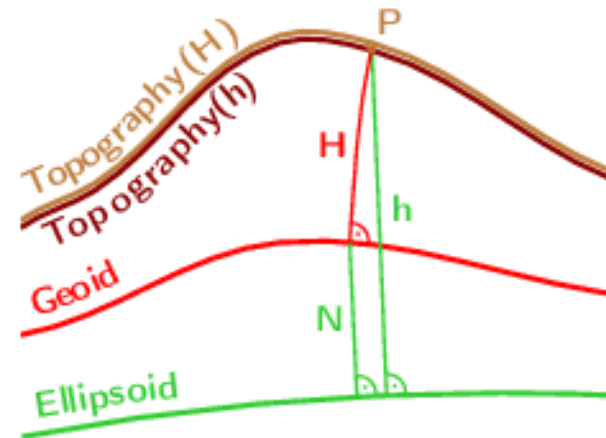
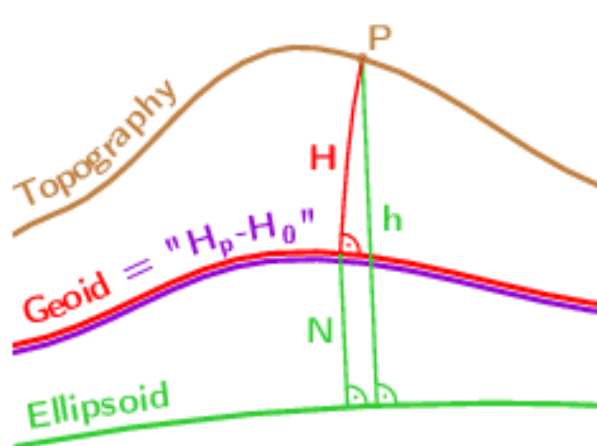
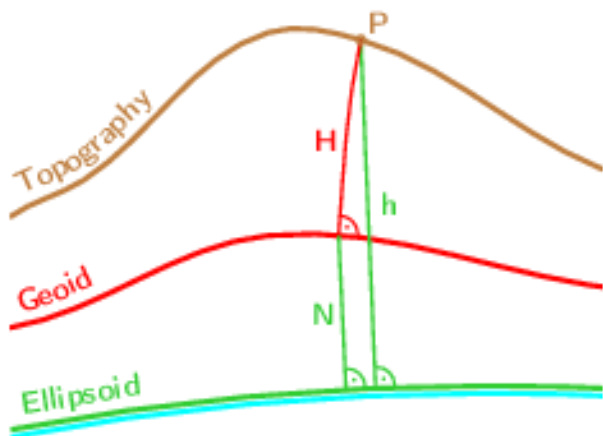
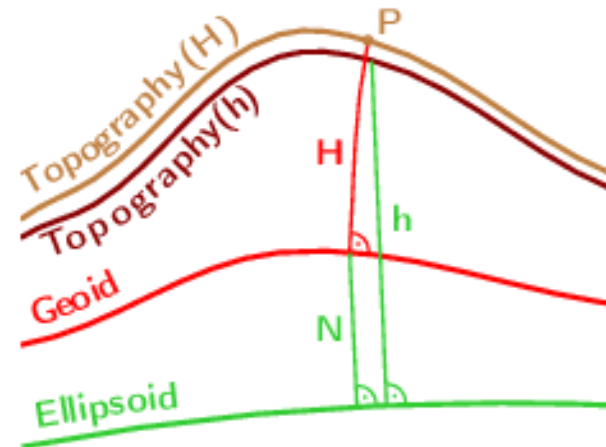
$h, N$  to get  $H$



$H, N$  to get  $h$

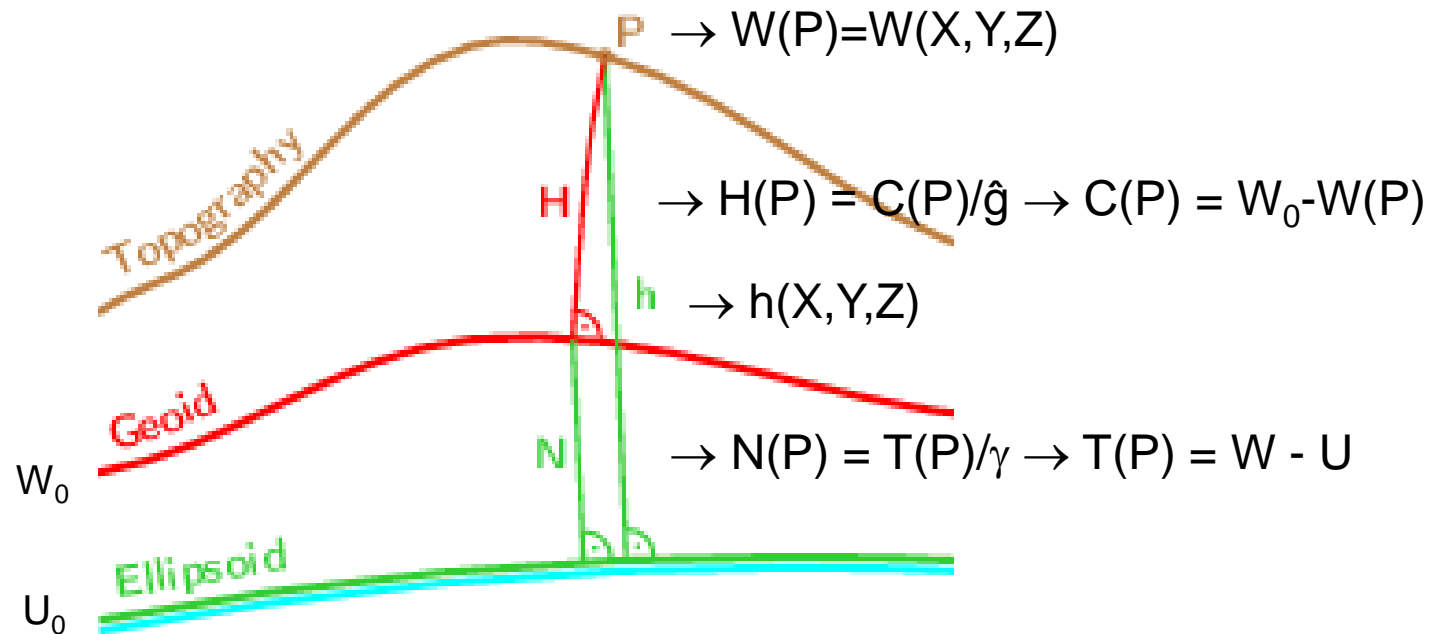


$h, H$  to get  $N$



worldwide and at the cm-level!

# Vertical coordinates in terms of potential



# Definition of the International Height Reference System (IHRIS)

## IAG Resolution No. 1, Prague, July 2015

- 1) Vertical coordinates are **potential differences** with respect to a **conventionally fixed  $W_0$**  value:

$$C_P = C(P) = W_0 - W(P) = -\Delta W(P)$$

$$W_0 = \text{const.} = 62\,636\,853.4 \text{ m}^2\text{s}^{-2}$$

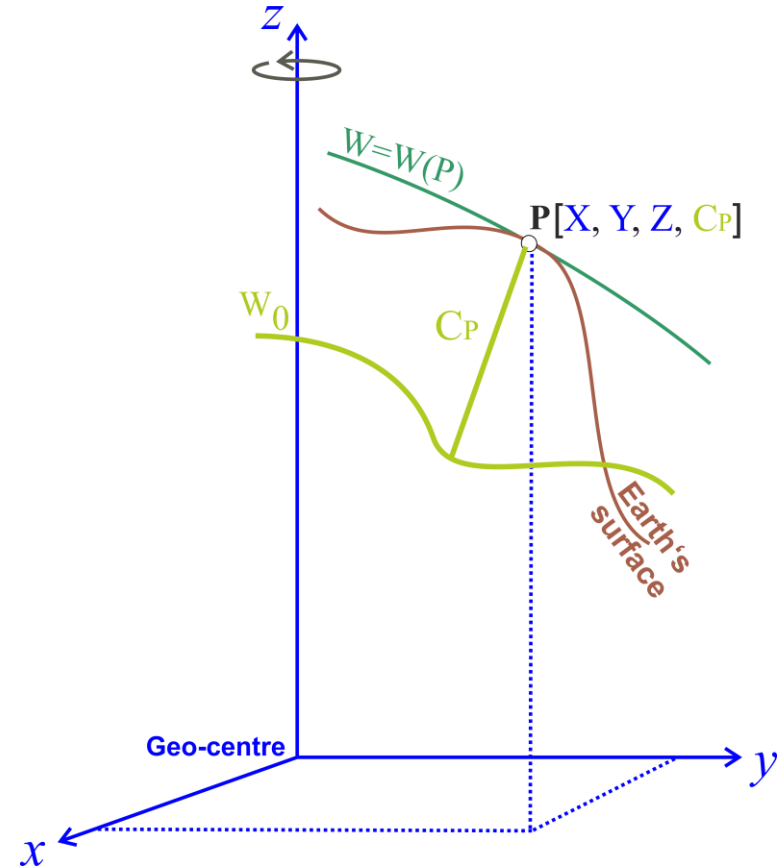
- 2) The position  $P$  is given in the ITRF

$$\mathbf{X}_P (X_P, Y_P, Z_P); \text{ i.e., } W(P) = W(\mathbf{X}_P)$$

- 3) The estimation of  $\mathbf{X}(P)$ ,  $W(P)$  (or  $C(P)$ ) includes their variation with time; i.e.,  $\dot{\mathbf{X}}(P)$ ,  $\dot{W}(P)$  (or  $\dot{C}(P)$ ).

- 4) Coordinates are given in **mean-tide system / mean (zero) crust**.

- 5) The unit of length is the **meter** and the unit of time is the **second (SI)**.



See: Ihde J. et al.: *Definition and proposed realisation of the International Height Reference System (IHRIS)*. *Surv Geophy* 38(3), 549-570, 10.1007/s10712-017-9409-3, 2017

# Realisation of the IHRS



A reference frame realises a reference system in two ways:

- physically, by a **solid materialisation of points** (or observing instruments),
- mathematically, by the **determination of coordinates** referring to that reference system.
- The coordinates of the points are computed from the measurements following the definition of the reference system.

During the last four years different actions have been conducted to

- Establish a global **reference network** for the IHRS realisation: the International Height Reference Frame (IHRF)
- Evaluate different **strategies for the determination of reference coordinates** at the reference stations
- Identify **required standards, conventions and procedures** needed to ensure consistency between the definition (IHRS) and the realisation (IHRF).

# Criteria for the IHRF reference network configuration



## 1) Hierarchy:

- A **global network** → worldwide distribution, including
- A **core network** → to ensure sustainability and long term stability
- **Regional and national densifications** → local accessibility

## 2) Collocated with:

- fundamental **geodetic observatories** → connection between  $\mathbf{X}$ ,  $W$ ,  $\mathbf{g}$  and time realisation (reference clocks) → **to support the GGRF**;
- **continuously operating reference stations** → to detect deformations of the reference frame (preference for ITRF and regional reference stations, like SIRGAS, EPN, APREF, etc.);
- **reference tide gauges and national vertical networks** → to facilitate the vertical datum unification;
- reference stations of the new **International Gravity Reference Frame - IGRF** (see IAG Resolution 2, Prague 2015).

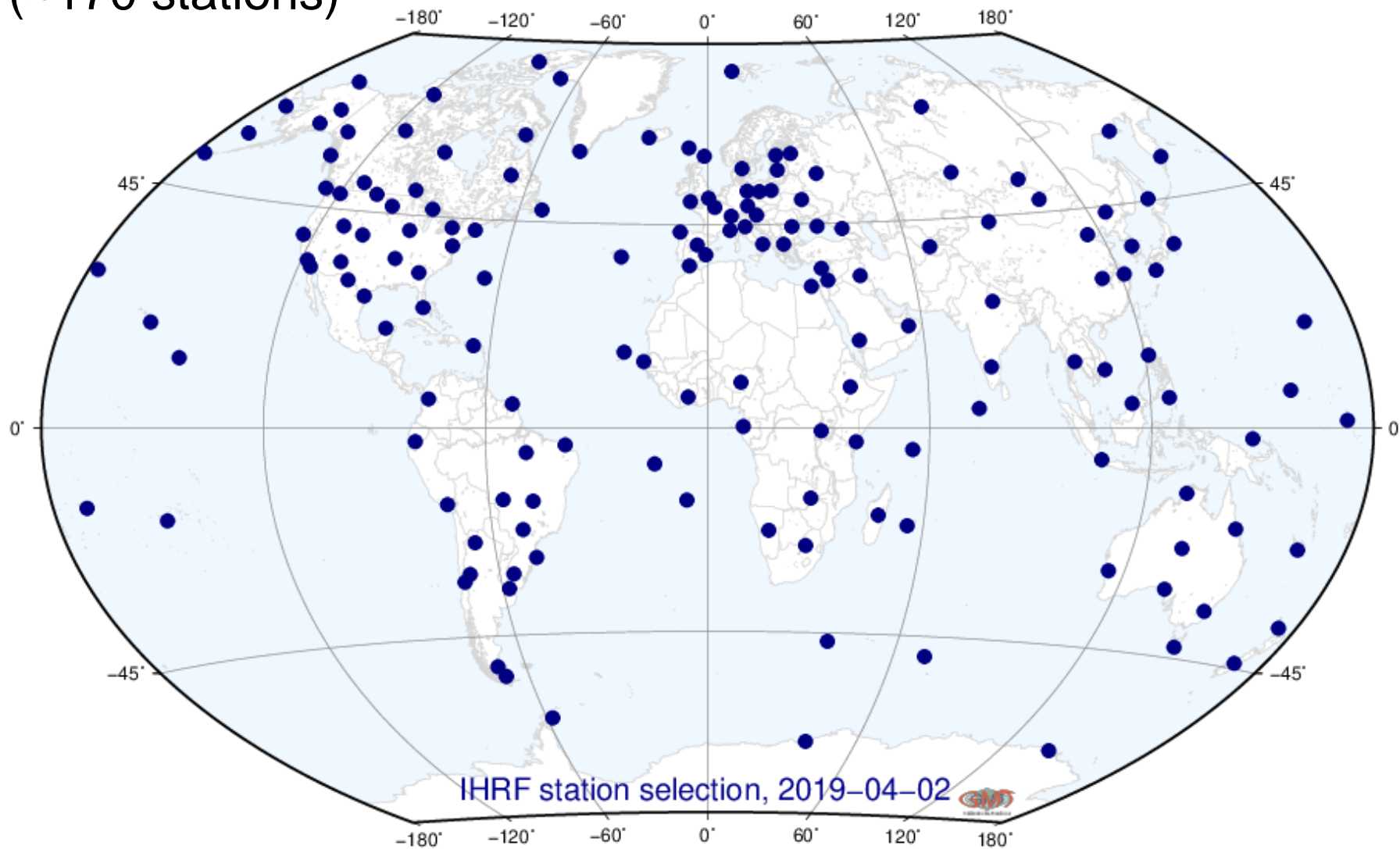
## 3) Main requirement: **availability of terrestrial gravity data around the IHRS reference stations for high-resolution gravity field modelling (i.e., precise estimation of $W$ ).**

# Station selection

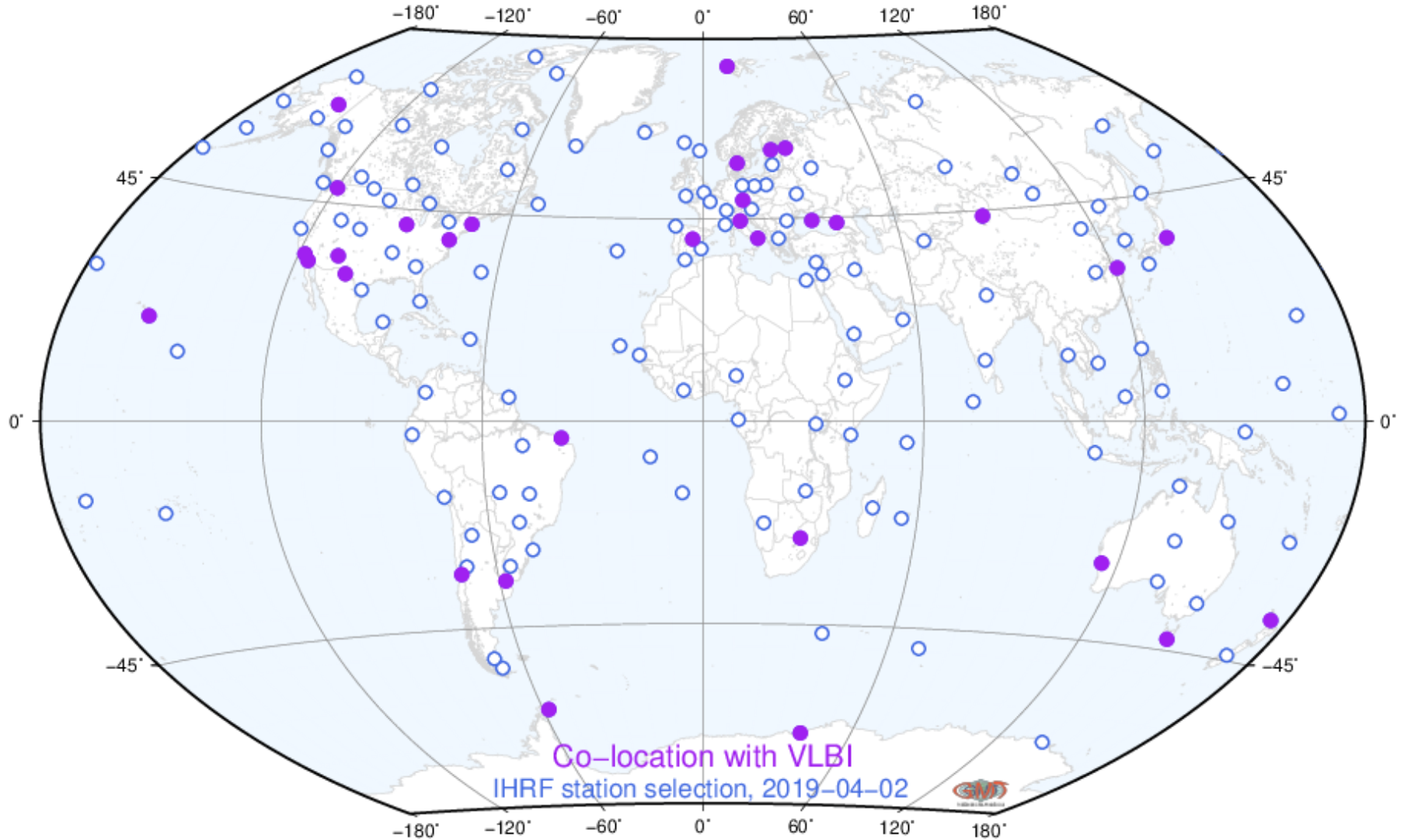


- 1) With the support of the [GGOS Bureau for Networks and Observations](#), a preliminary selection based on [VLBI](#), [SLR](#) and [DORIS](#) reference sites co-located with [GNSS](#) was prepared (Oct 2016).
- 2) Based on these preliminary selection, national/regional experts were asked to
  - evaluate whether these sites are suitable to be included in the IHRF: Are gravity data around these sites available? If not, is it possible to survey gravity around them?
  - propose additional geodetic sites to improve the density and distribution of the IHRF stations in their regions/countries
- 3) With support of the IAG JWG 2.1.1: [Establishment of a global absolute gravity reference system](#) (chair: H. Wziontek), further stations co-located with absolute gravity stations were identified.
- 4) A first proposal for the IHRF reference network was ready in Apr 2017.
- 5) Since that time some new stations have been added, others have been decommissioned.
- 6) It is expected that this network is extended by means of regional/national densifications.

# First proposal for the IHRF reference network (~170 stations)

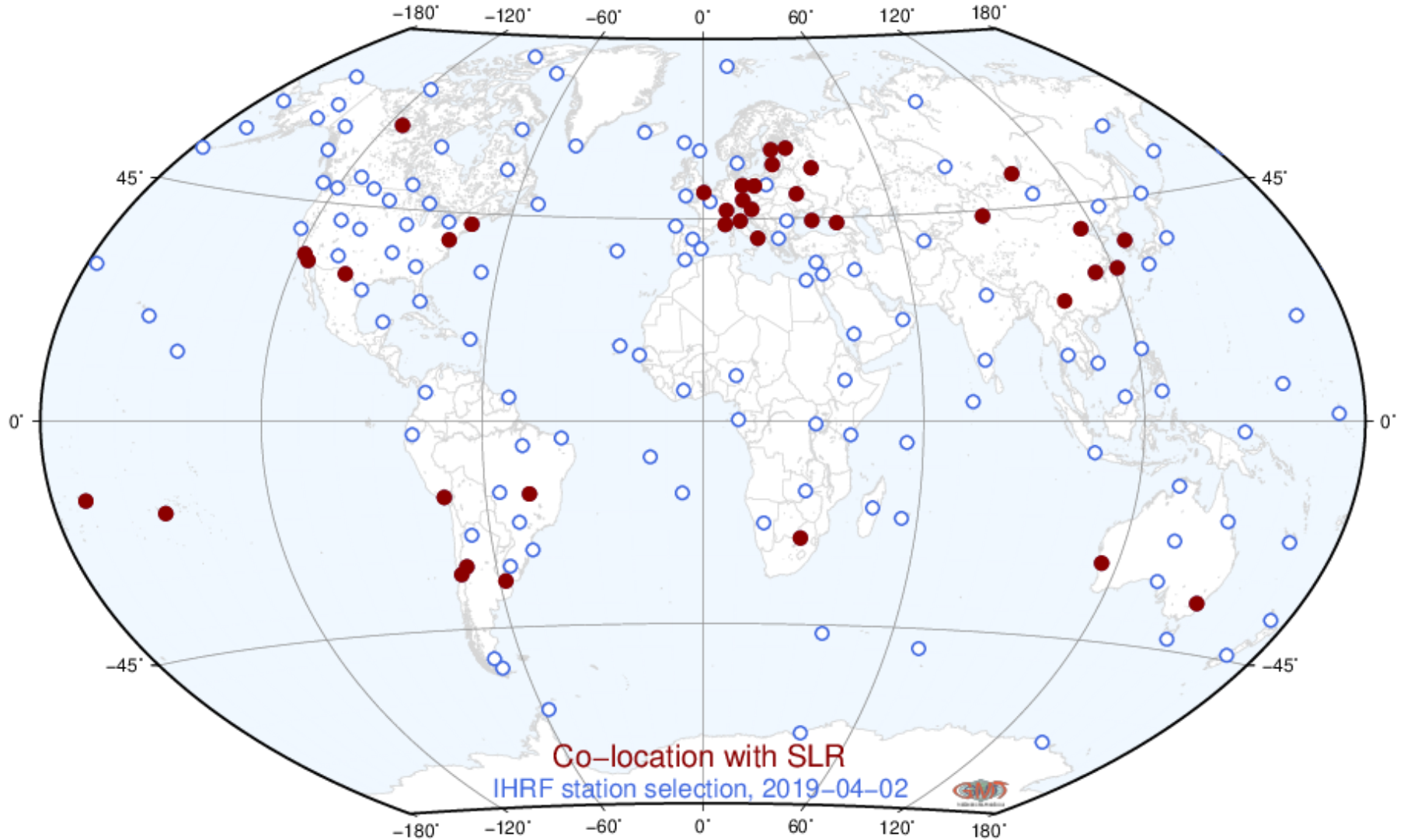


# Co-location with VLBI (30 sites)

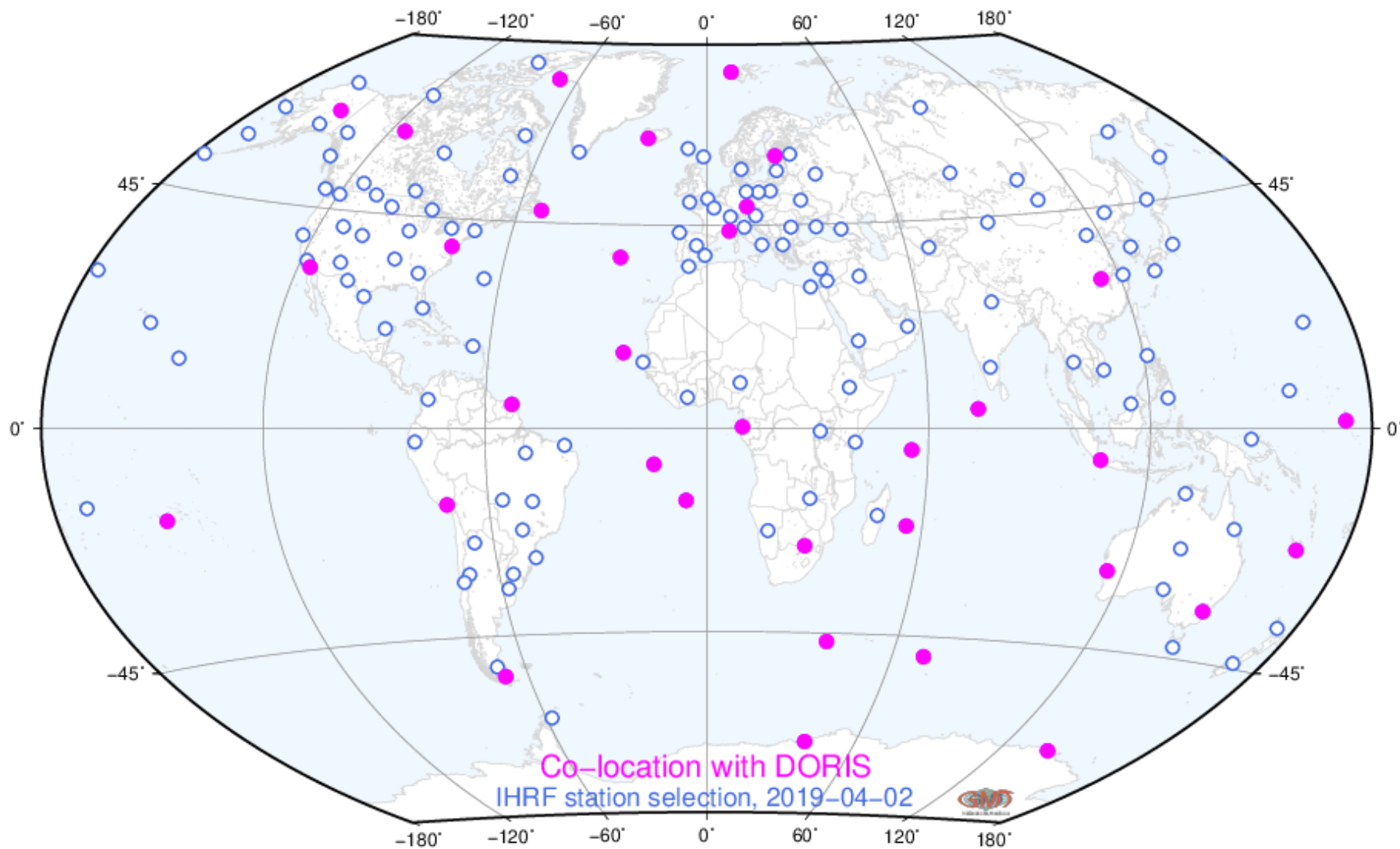




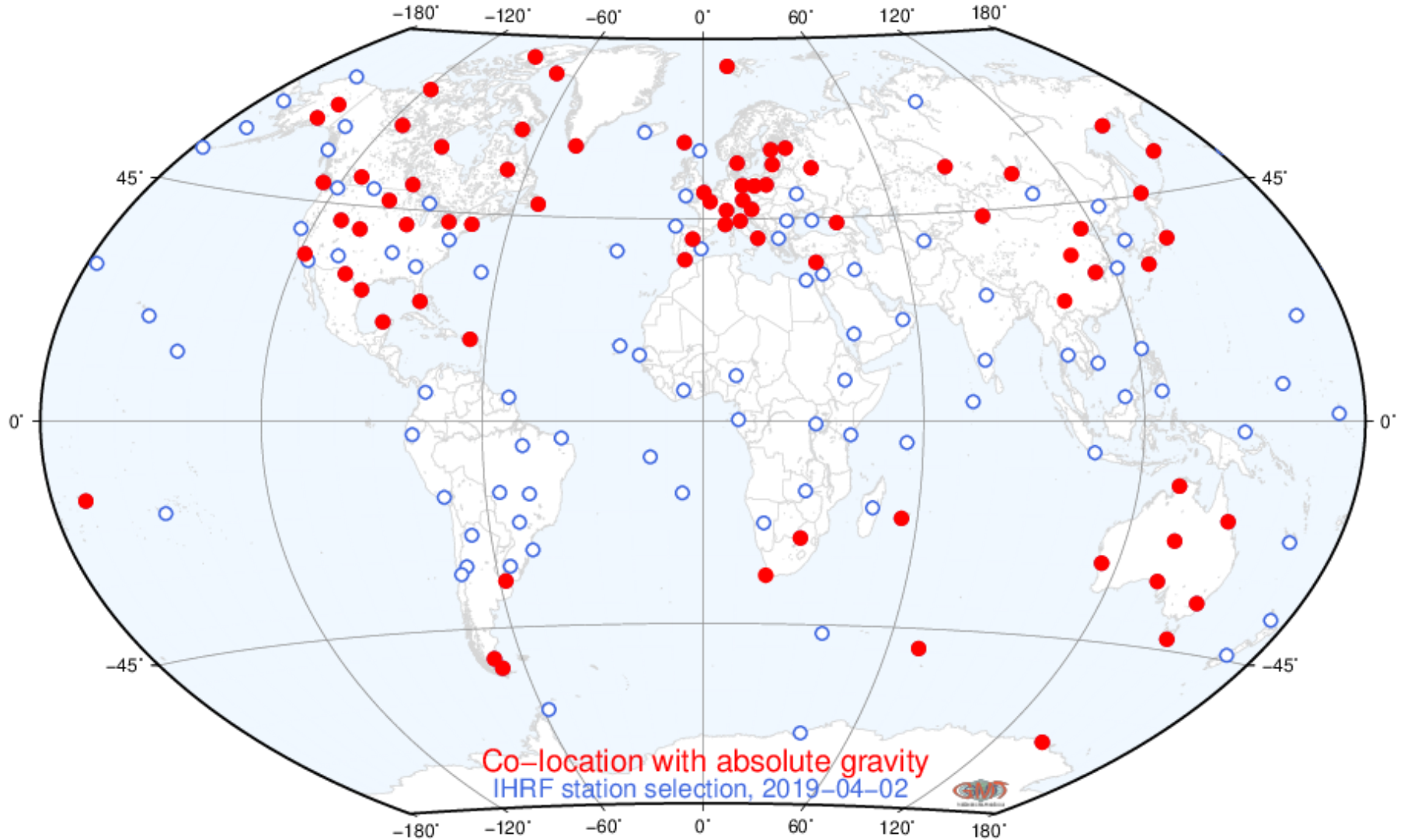
# Co-location with SLR (40 sites)



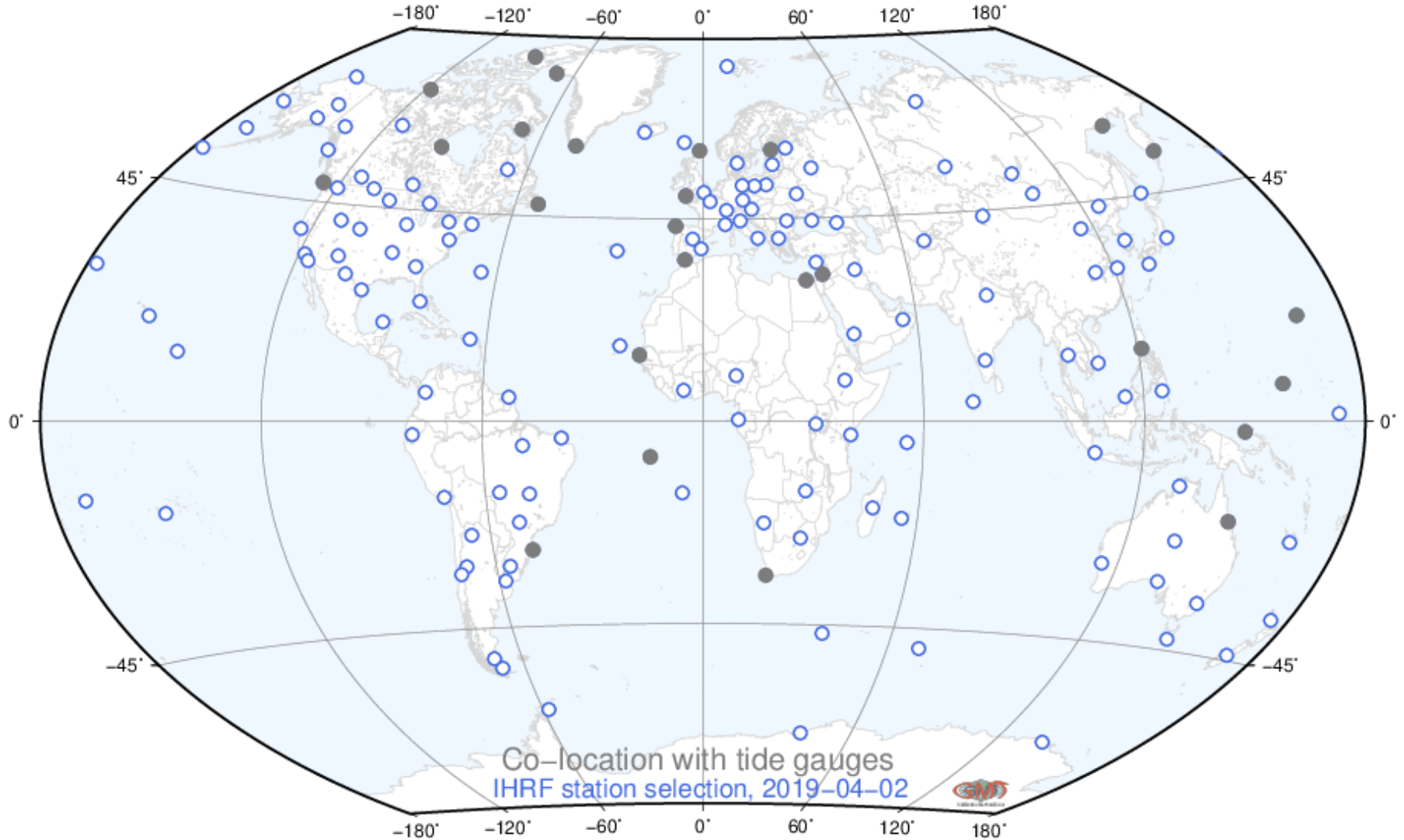
# Co-location with DORIS (35 sites)



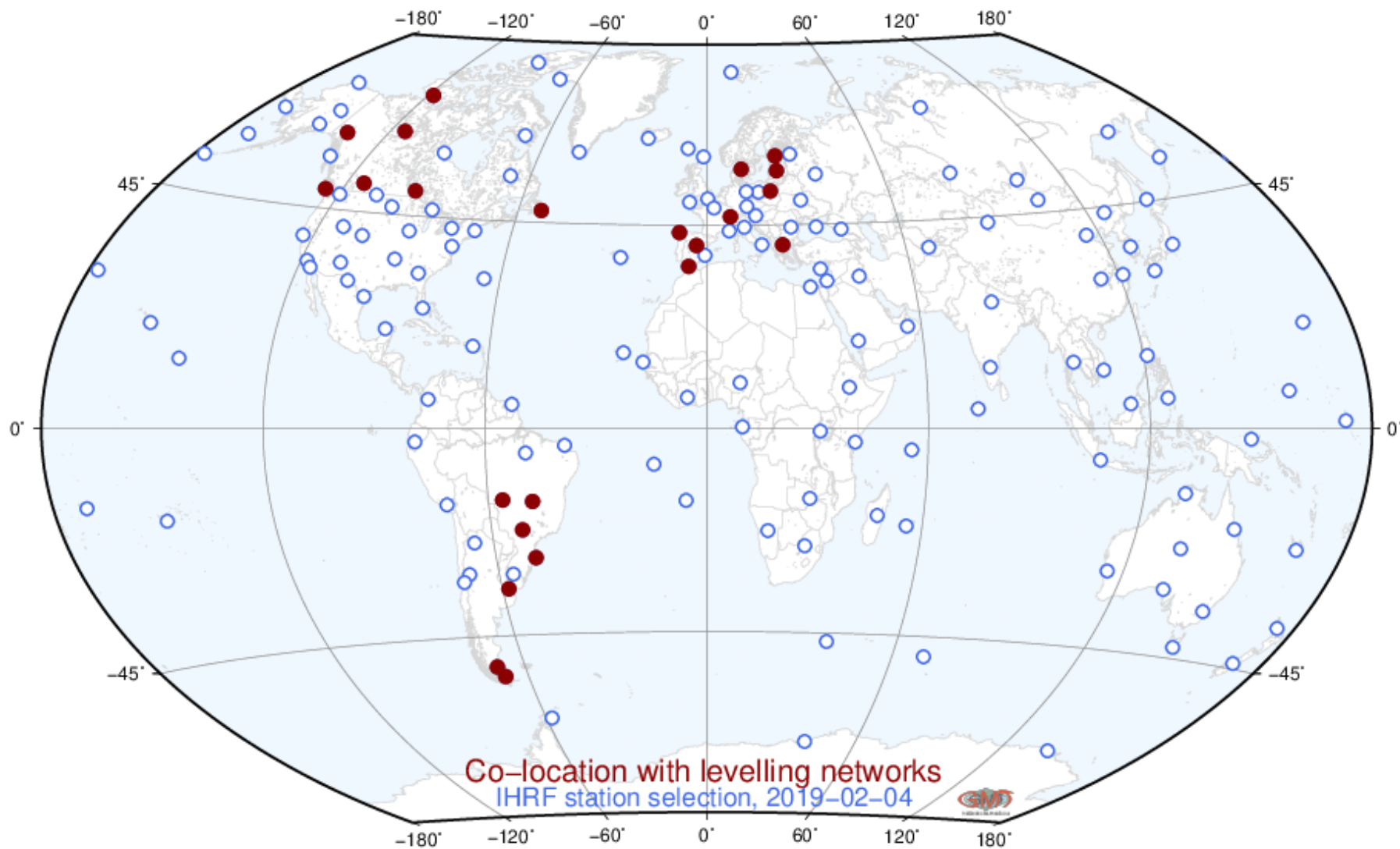
# Co-location with absolute gravity (77 sites)



# Co-location with tide gauges (26 sites)



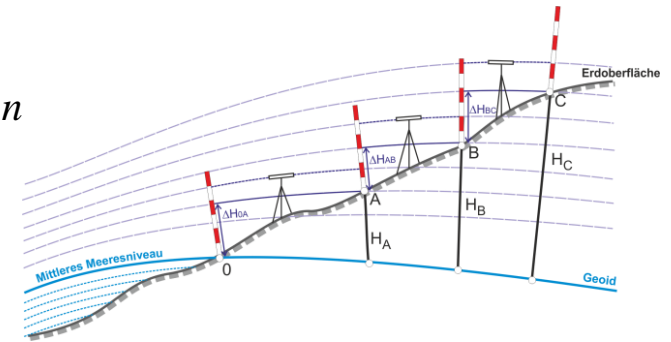
# Co-location with levelling networks (23 sites)



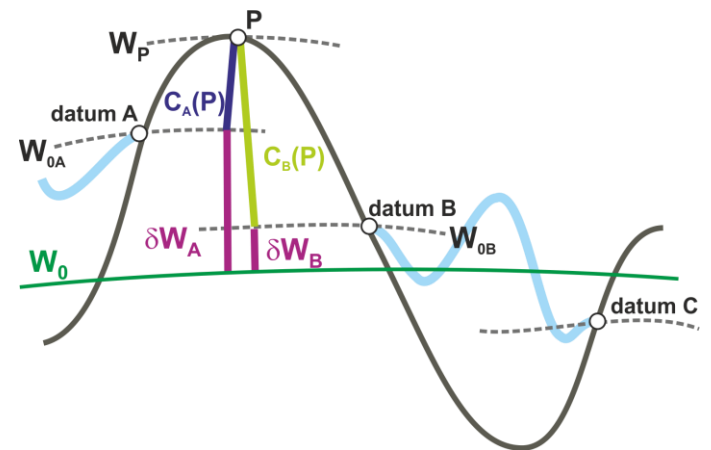
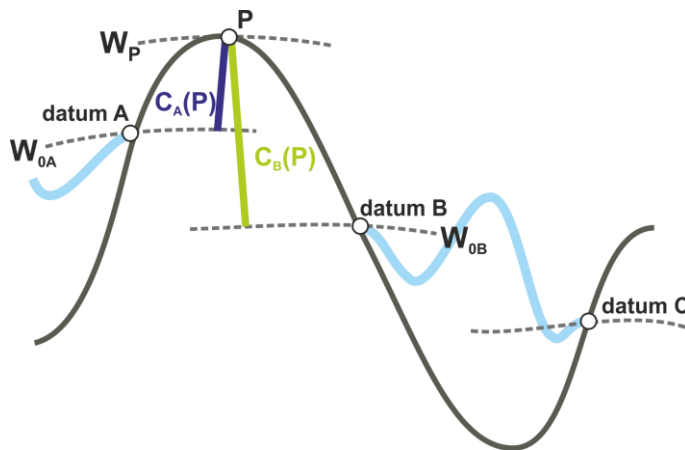
- 1) The IHRS/IHRF is the combination of a **geometric component** given by the coordinate vector  $\mathbf{X}$  in the ITRS/IHRF and a **physical component** given by the determination of potential values  $W$  at  $\mathbf{X}$ .
- 2) The determination of  $\mathbf{X}$  follows the **IERS Conventions**. There is not something similar to the IERS Conventions for the determination of  $W$ .
- 3) Current target accuracy for vertical coordinates:
  - Accuracy of the **geoid** (geometry of any equipotential surface)
    - Static geoid:  $\pm 1$  mm, spatial resolution: 10 km.
    - Time-dependent geoid:  $\pm 1$  mm, spatial res. 50 km, temporal res. 10 days
  - Accuracy of the **ITRF coordinates**:
    - Positions:  $\pm 1$  mm horizontal,  $\pm 3$  mm vertical.
    - Velocities:  $\pm 0.1$  mm/a horizontal,  $\pm 0.3$  mm/a vertical.
  - Inferred (expected) accuracy for  $W_p$ :
    - Positions:  $\approx \pm 3 \times 10^{-2} \text{ m}^2\text{s}^{-2}$  (about  $\pm 3$  mm).
    - Velocities:  $\approx \pm 3 \times 10^{-3} \text{ m}^2\text{s}^{-2}$  (about  $\pm 0.3$  mm/a).
- 4) For the moment, our goal is  $\pm 1 \times 10^{-1} \text{ m}^2\text{s}^{-2}$  (about 1 cm)
- 5) The IHRS/IHRF coordinates include the determination of time variations. For the moment, we consider **static coordinates only**.

## 1) Geopotential numbers inferred from levelling and gravity reductions:

$$W_P = (W_0^{local} + \delta W) - C_P; \quad \delta W = W_0^{IHRF} - W_0^{local}; \quad C_P = \int_0^P g \, dn$$



- Refer to local vertical datums with unknown potential value  $W_{0,local} = ?$
- To determine  $W_P$ , it is necessary to estimate the level difference between the global  $W_0$  and the local  $W_{0,local} \rightarrow \delta W = W_0 - W_{0i}$

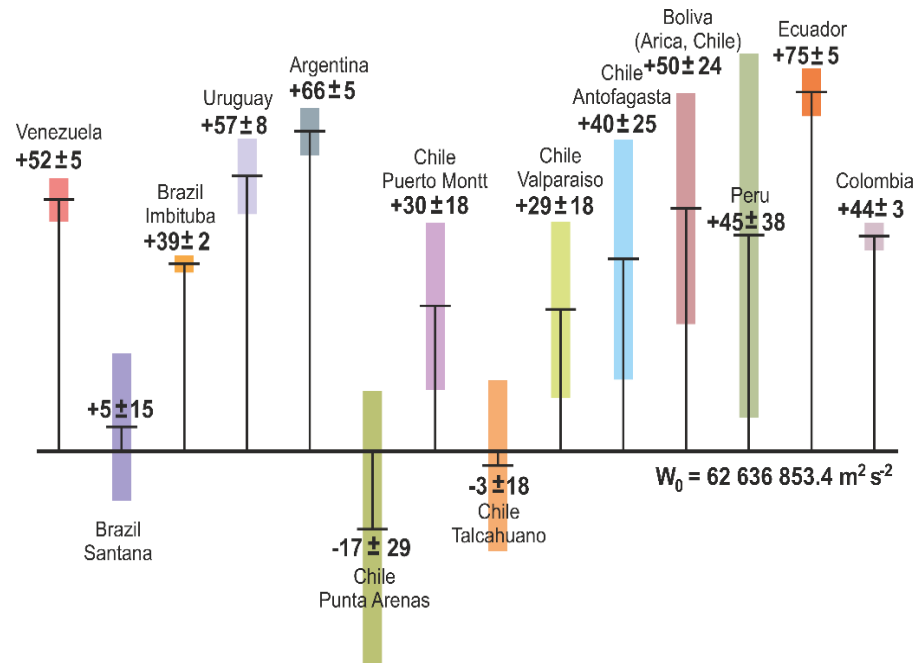


# Possibilities for the determination of potential values

## 1) Geopotential numbers inferred from levelling and gravity reductions:

- Example:  $\delta W$  (in cm) for the South American height systems w.r.t. the IHRs  $W_0$  value.
- Reliability depends on the **limitations of the existing height systems**, in particular
  - the strong accumulation of **systematic errors in levelling**, and
  - the impossibility of referring the levelled heights to a **specific epoch**

▪ This approach is useful for the **transformation of the existing height systems to the IHRs**, but it may be unsuitable for the precise realisation of the IHRs.

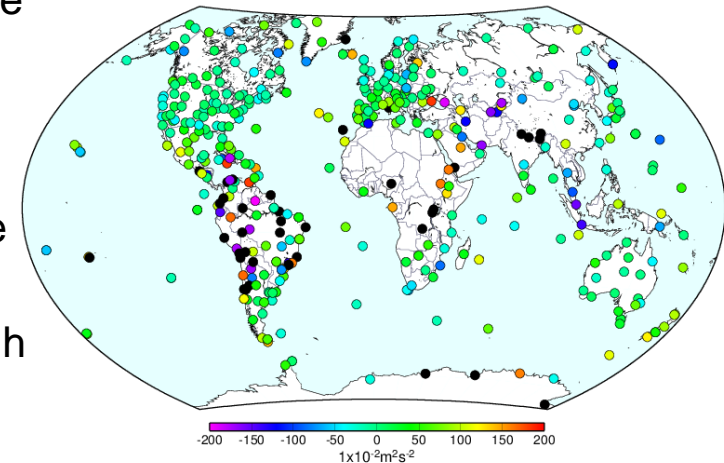




- 2) **Global Gravity Models of high degree (GGM-HD)** like the EGM2008 model (Pavlis et al., 2012, 2013) or the EIGEN-C series (e.g., Förste et al., 2012; 2014)

$$W(X, Y, Z) = \frac{GM}{r} \left[ 1 + \sum_{n=1}^{\infty} \left( \frac{a}{r} \right)^n \sum_{m=0}^n [C_{nm} \cos m\lambda + S_{nm} \sin m\lambda] P_{nm}(\cos\theta) \right] + \frac{1}{2} \omega^2 r^2 \cos(90^\circ - \theta)$$

- Expected accuracy (Rummel et al., 2014)
  - well surveyed regions:  $\pm 0.4 \text{ m}^2\text{s}^{-2}$  to  $\pm 0.6 \text{ m}^2\text{s}^{-2}$  (equivalent to  $\pm 4 \text{ cm}$  to  $\pm 6 \text{ cm}$ )
  - sparsely surveyed regions: from  $\pm 2 \text{ m}^2\text{s}^{-2}$  ...  $\pm 4 \text{ m}^2\text{s}^{-2}$  ( $\pm 20 \text{ cm}$  to  $\pm 40 \text{ cm}$ ) to  $\pm 10 \text{ m}^2\text{s}^{-2}$  ( $\pm 1 \text{ m}$ )
- Differences between the  $W_p$  values derived from EGM2008 (Pavlis et al. 2008) and EIGEN6C4 (Förste et al. 2014), both at  $n=2190$ 
  - Differences larger than  $\pm 200 \times 10^{-2} \text{ m}^2\text{s}^{-2}$  ( $\approx \pm 2 \text{ m}$ )
  - Desired accuracy for  $W_p$ :  $\pm 0.03 \text{ m}^2\text{s}^{-2}$  ( $\approx \pm 3 \text{ mm}$ )
- This approach represents the “ideal way” to estimate potential values and hopefully, we will get a better accuracy in the next future. Ongoing studies with high expectation of improvement:
  - Combination of GGM with gravity effects of global topography
  - EGM2020
  - However, **terrestrial gravity data is further required!**



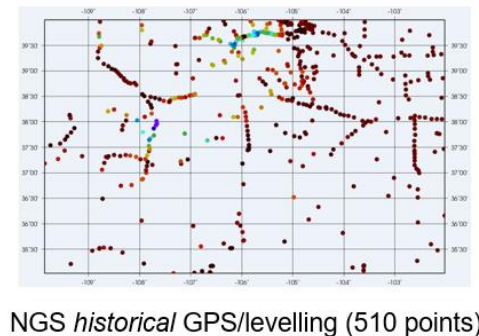
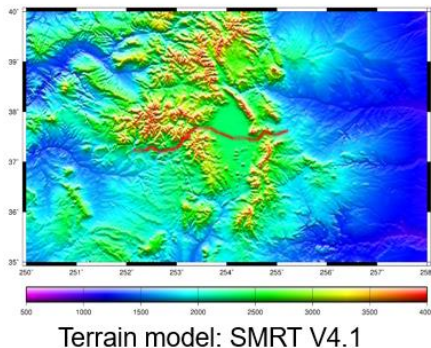
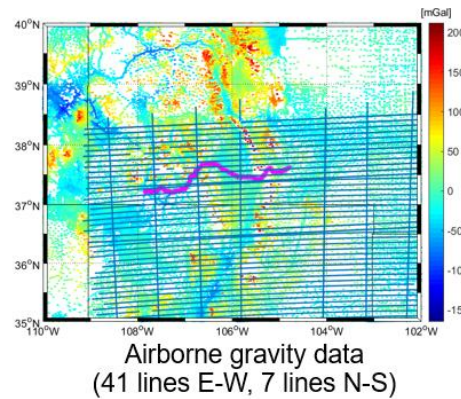
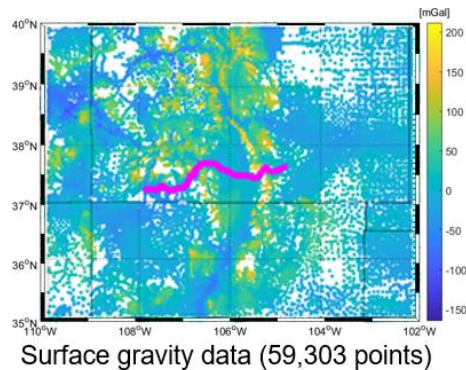
## 3) Disturbing potential

$$W_P = U_P + T_P ; \quad T_P = T_{P,satellite-only} + T_{P,residual} + T_{P,terrain}$$

- GGM based on SLR, GRACE and GOCE are **very precise** ( $\pm 1 \dots \pm 2$  cm @ 100 km)
- Mean omission error globally:  $\approx \pm 45$  cm
- Goal is to **reduce these  $\pm 45$  cm to  $\pm 1$  cm** (only possible using terrestrial gravity data and considering topographic effects)
- The potential values realising the ITRS coordinates must be determined at the reference stations; i.e., at the Earth's surface and not at the geoid
- The determination of  $T_P$  demands a series of approximations, which influence the results; i.e., **different methodologies produce different potential values**

# Comparison of computation methods

Colorado experiment: to compute geoid, quasi-geoid and potential values using exactly the same input data, a set of basic standards, and the own methodologies (software) of colleagues involved in the gravity field modelling.



- Initiated in July 2017
- Data provided by US NGS
- Standards prepared by L Sánchez, J Ågren, J Huang, YM Wang, R Forsberg
- Three computations (two iterations) finished in June 2019
- Fifteen (final) contributing solutions
- Special Issue of the Journal of Geodesy with computation methods and comparison of geoid and quasi-geoid models (in preparation).

# Colorado experiment: contributing solutions



Faculty of Engineering, Minia University, [Egypt](#)



İstanbul Teknik Üniversitesi, Istanbul, [Turkey](#)



Department of Geodesy and Surveying, Aristotle University of Thessaloniki, Thessaloniki, [Greece](#)



National Geodetic Survey, [USA](#)



Natural Resources Canada, [Canada](#)

LANTMÄTERIET

Lantmäteriet, Swedish mapping, cadastral and land registration authority, [Sweden](#)



School of Earth and Planetary Sciences and The Institute for Geoscience Research, Curtin University, [Australia](#)



Escola Politécnica, Universidade de São Paulo; Centro de Estudos de Geodesia, [Brazil](#)



Deutsches Geodätisches Forschungsinstitut, Technische Universität München, [Germany](#)



Ingenieurinstitut für Astronomische und Physikalische Geodäsie, Technische Universität München, [Germany](#)



Chinese Academy of Surveying and Mapping, [China](#)



POLITECNICO MILANO 1863

Politecnico de Milano, [Italy](#)



Faculty of Geodesy, University of Zagreb, [Croatia](#) - Research Institute of Geodesy, Topography and Cartography, [Czech Republic](#)



National Space Institute, Technical University of Denmark, [Denmark](#)



Geography and Crustal Dynamics Research Center, Geospatial Information Authority of Japan, [Japan](#)

# Colorado experiment: summary of approaches and models



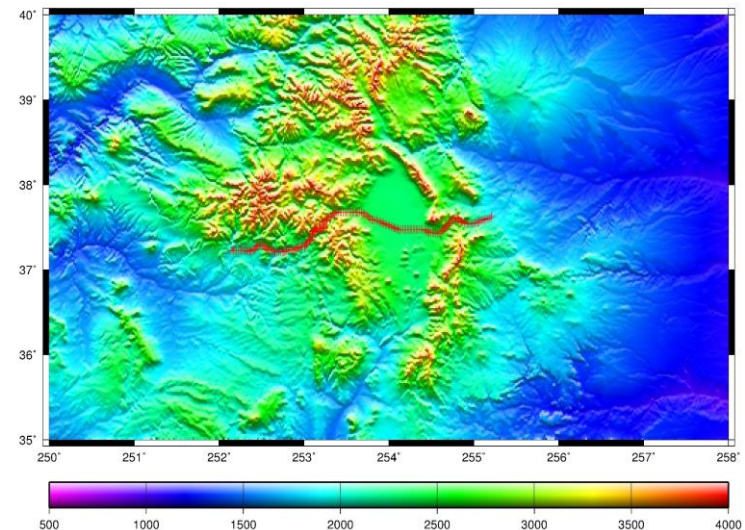
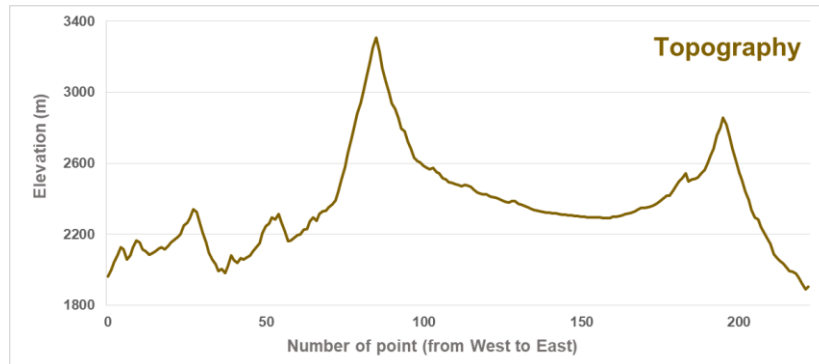
- Least squares modification of Stokes' formula with additive corrections (2)
- Least squares modification of Stokes' formula with additive corrections and biased Stokes' kernel modification
- Stokes' formula with Wong-Gore modification and 1D-FFT (2)
- Spherical radial basis functions
- Least squares collocation
- Fast collocation based on gravity gridded data
- Degree weighted Stokes' integral
- Modified degree-banded Stokes' kernel (2)
- Spherical FFT with modified Wong-Core Stokes' kernel
- UNB Stokes-Helmert scheme
- UNB Stokes-Helmert scheme with hybrid-Meissl-Molodensky modified spheroidal Stokes' kernel
- NGS Molodensky approach, Spherical Harmonics Analysis (SHA)
- GGMs: GOCO05s, XGM2016, XGM2018, xGEOID17B, EIGEN-6C4, EGM2008
- Topographic effects based on SRTM V4.1, EARTH2014, COLH19M05, ERTM2160
- 12 solutions based on height anomalies, 3 solutions based on geoid undulations

# Colorado experiment: comparison of potential values

- 1) The comparison is carried out at 223 GSVS17 marks (Geoid Slope Validation Survey 2017) selected by NGS
- 2) Participants in the experiment got  $\varphi$ ,  $\lambda$ ,  $h$ ; levelling is not available (yet)
- 3) The potential values provided by the different solutions are converted to **geopotential numbers** with respect to the IHRs  $W_0$  value

$$C(P) = W_0 - W(P) \quad ; \quad W_0 = 62\,636\,853.4 \text{ m}^2\text{s}^{-2}$$

- 4) and further transformed to **normal heights** (to see the differences in meters):  $H^*(P) = C(P)/\gamma(P)$



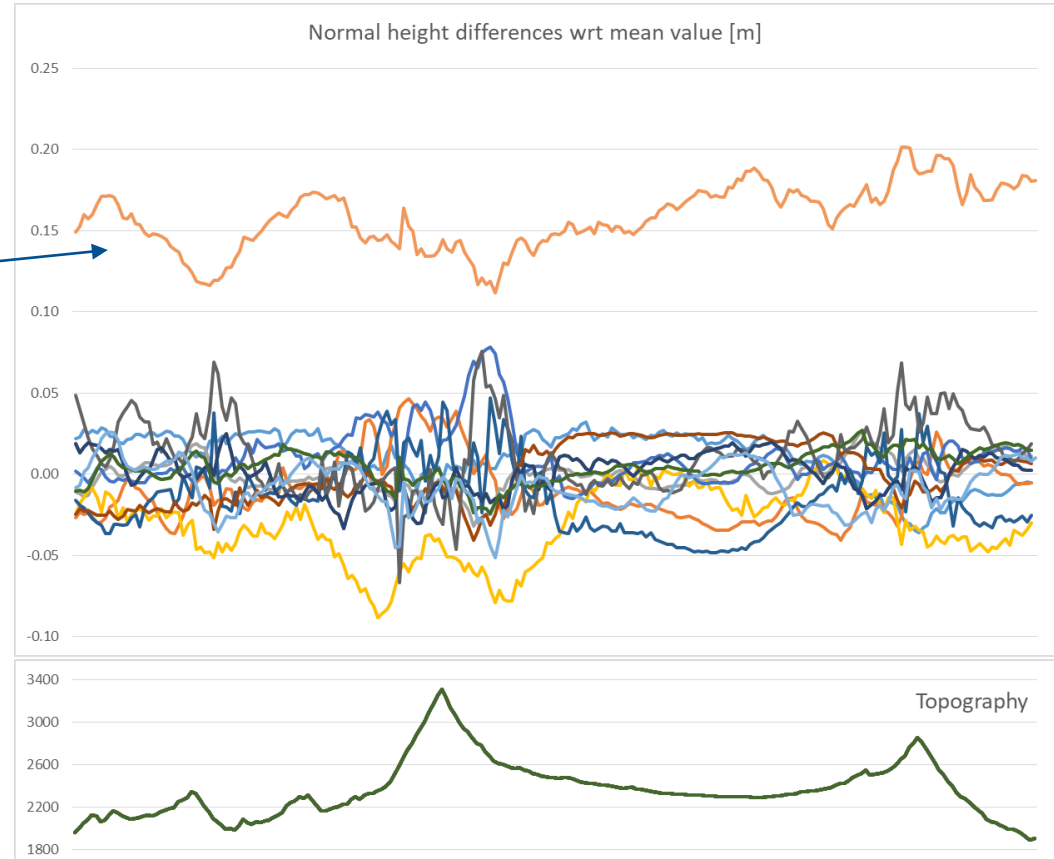
# Colorado experiment: comparison of potential values

Normal height difference [cm]  
(individual contribution – mean)

Outlier 1

Mean :  $15.7 \pm 1.9$  cm

Range: 8.9 cm (11.2 ... 20.0 cm)



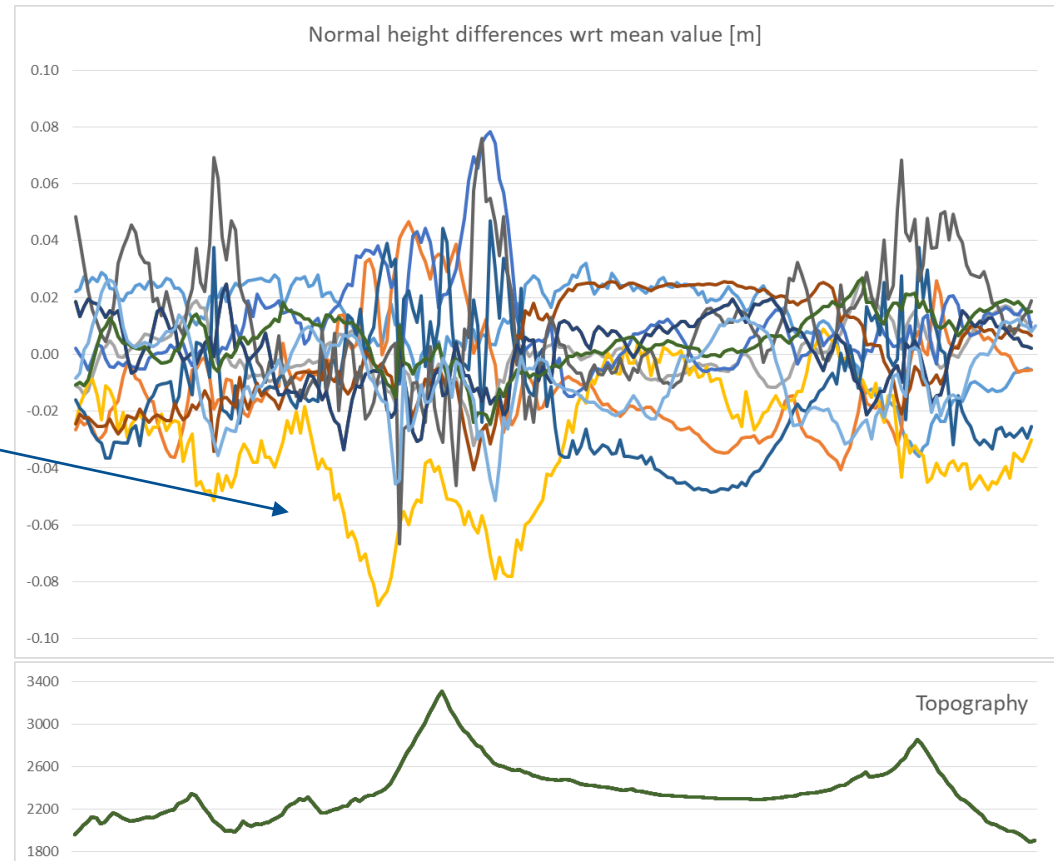
Zero-degree term: 17.85 cm

$$\zeta_0 = \frac{(GM_{GGM} - GM_{GRS80})}{r_P \cdot \gamma_Q} - \frac{\Delta W_0}{\gamma_Q}$$

# Colorado experiment: comparison of potential values

Normal height difference [cm]  
(individual contribution – mean)

Outlier 2  
Mean :  $-3.2 \pm 2.1$  cm  
Range: 9.3 cm (-8.7 ... 0.6 cm)

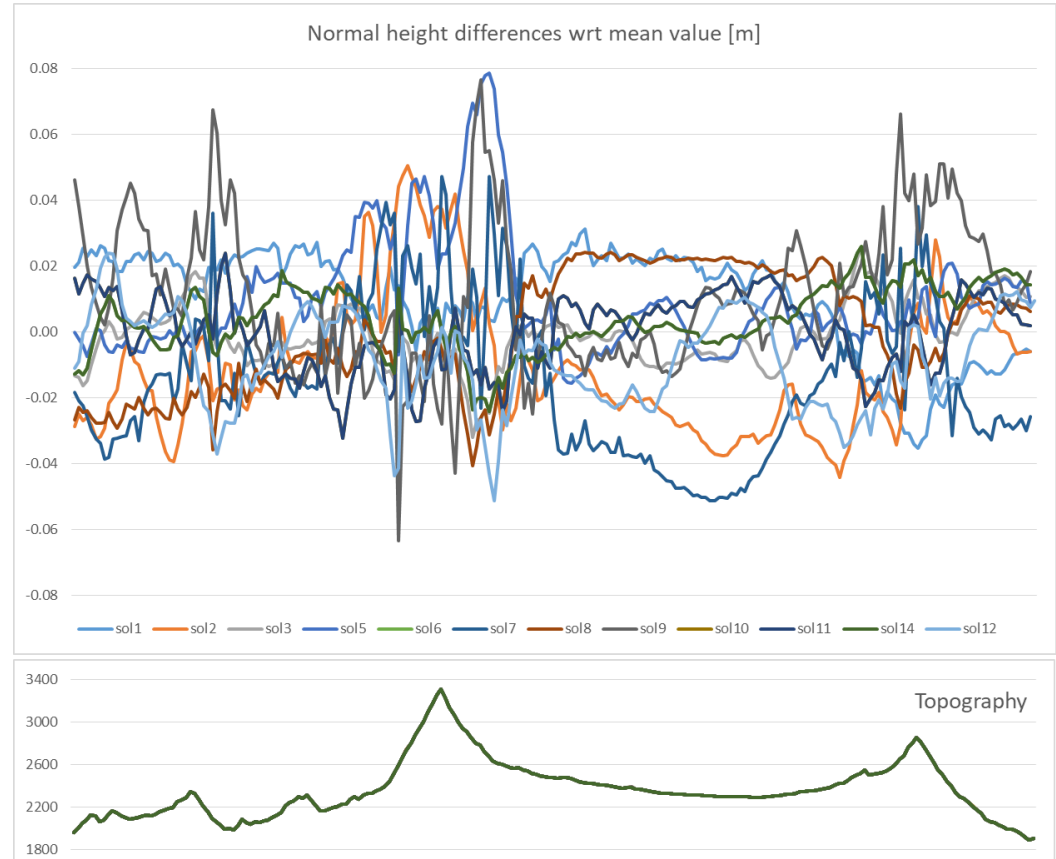




# Colorado experiment: comparison of potential values

Normal height difference [cm]  
(individual contribution – mean)

	Mean	±	STD	Range
sol1	1.0	±	1.6	6.6 ( -3.5 ... 3.1 )
sol2	-1.0	±	2.1	9.5 ( -4.4 ... 5.1 )
sol3	-0.1	±	1.0	5.8 ( -3.2 ... 2.6 )
sol5	1.0	±	1.8	9.4 ( -1.6 ... 7.9 )
sol6	0.4	±	1.0	5.3 ( -2.7 ... 2.6 )
sol7	-1.4	±	2.3	9.9 ( -5.1 ... 4.7 )
sol8	0.0	±	1.8	6.5 ( -4.1 ... 2.4 )
sol9	1.1	±	2.2	14.0 ( -6.3 ... 7.7 )
sol10	0.0	±	1.2	7.5 ( -3.2 ... 4.3 )
sol11	0.0	±	1.1	5.6 ( -3.2 ... 2.4 )
sol12	-0.9	±	1.4	7.5 ( -5.1 ... 2.4 )
sol14	0.4	±	1.0	5.0 ( -2.4 ... 2.6 )



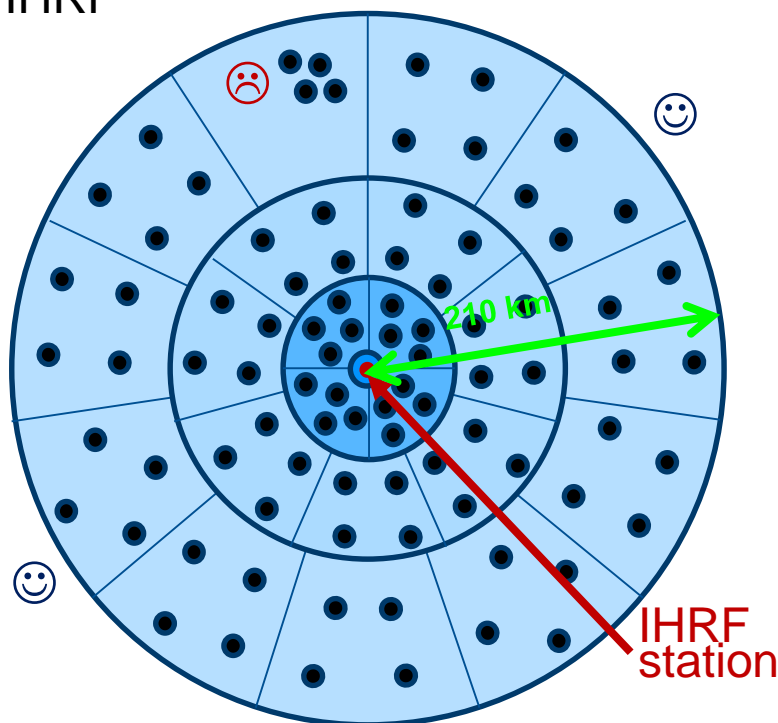
# Learnings from the Colorado experiment



- 1) Validation of gravity field (geoid) modelling additional to GNSS/levelling
- 2) Twelve(!) solutions agree within 1 cm to 2 cm in terms of standard deviation with respect to the mean value
- 3) We are waiting for the levelling results along the test profile to make comparisons with independent data
- 4) Discrepancies between the different solutions are highly correlated with the topography
  - Handling of terrain gravity effects (model and strategy)
- 5) Difficulties reported by the colleagues contributing to the experiment
  - Processing of the airborne gravity data
  - Handling of the zero-degree term
- 6) A major confusion is the reference ellipsoid: which should be used GRS80 or WGS84?
  - Are we needing a new reference ellipsoid?

# Learnings of the Colorado experiment

- The GGM should be based at least on the combination of SLR, GRACE and GOCE data ( $n \geq 200$ )
- To get an accuracy of about  $\pm 1$  cm in the (quasi-)geoid, observed gravity values are required with a mean spatial resolution of about 4 km
- The availability of these data is a main criterion to select reference stations for the IHRF



**Template according to the gravity effect on the geoid ( $\Delta g = 1 \cdot 10^{-6} \text{ ms}^{-2} \rightarrow 1 \text{ mm}$ )**

Distance	Compartments	# of points <b>flat/mountain</b>
10 km	1	<b>4/8</b>
50 km	4	<b>20/30</b>
110 km	7	<b>30/45</b>
210 km	11	<b>50/75</b>
Sum	23	<b>104/158</b>

### Resolution 3: Establishment of the International Height Reference Frame (IHRF)

The International Association of Geodesy,

*Considering,*

- The IAG Resolution for the Definition and Realization of an International Height Reference System (IHRF) released at the 26th IUGG General Assembly in July 2015;

*Acknowledging,*

- The achievements of
  - GGOS Focus Area "Unified Height System" and its JWG 0.1.2 "Strategy for the Realization of the International Height Reference System (IHRF)",
  - IAG JWG 2.2.2 "The 1 cm geoid experiment",
  - IAG SC 2.2 "Methodology for geoid and physical height systems",
  - ICCT JSG 0.15 "Regional geoid/quasi-geoid modelling - Theoretical framework for the sub-centimetre accuracy";
- in realizing this resolution;

*Noting,*

- The need of an operational infrastructure to ensure the determination, maintenance and availability of an International Height Reference Frame (IHRF) in the long-term basis;

*Urges,*

- All countries to engage with the IAG and concerned components, in particular the International Gravity Field Service (IGFS), in order to promote and support the implementation of the IHRF by
  - Installing IHRF reference stations at national level,
  - Conducting the necessary gravimetric surveys to guarantee the precise determination of potential values,
  - Making data available open access,
  - Contributing to the development of analysis strategies to improve the estimation of reference coordinates and modelling of the Earth's gravity field,
  - Describing, archiving and providing geodetic products associated to the IHRF.

- 1) To compute a **first static solution for the IHRF** to evaluate the achievable accuracy under the present conditions (data availability, computation methods, etc.) and to identify key actions to improve the determination of the IHRS/IHRF coordinates.
- 2) To investigate the determination of **potential changes** with time  $\dot{W}$ .
- 3) To extend the realisation of the **IHRF to marine areas**.
- 4) To explore the possibilities to establish an '**IHRF element**' within the International Gravity Field Service (IGFS) to ensure the maintenance and availability of the IHRF:
  - Regular updates of the IHRF<sub>yyyy</sub> to take account for:
    - new stations;
    - coordinate changes with time  $\dot{\mathbf{X}}$ ,  $\dot{W}$ ;
    - improvements in the estimation of  $\mathbf{X}$  and  $W$  (more observations, better standards, better models, better computation algorithms, etc.)

# Participation of Latin America in the implementation of the IHRS/IHRF



## 1) Establishment of IHRS stations

- To select some (1 to 5) continuously operating SIRGAS reference stations in each country (well distributed and materialized by a monument on the ground; stations on the top of buildings are not welcome).
- To survey gravity data around the selected SIRGAS reference stations (about 150 gravity points well distributed around each station up to a distance of about 200 km).
- Coordinates of gravity points determined with GNSS positioning ( $\pm 2$  cm).
- It is desirable that the gravity surveys refer to absolute gravity stations.

## 2) Integration of the existing Latin American height systems into the IHRS/IHRF

- First order levelling (with gravity data) of SIRGAS reference stations (optimal if IHRF stations are levelled).
- Reference tide gauges connected to SIRGAS.
- Combination of ellipsoidal heights, levelling-based physical heights, tide gauge registrations, satellite altimetry observations and height-resolution gravity field modelling.

## 3) Latin American countries should take advantage of the SIRGAS-WG3 activities:

- Capacity building and software for the processing of gravity data
- Capacity building and software for the adjustment of levelling networks and computation of geopotential numbers
- Until now: Rio (2012), La Paz (2014), Curitiba (2015), Quito (2016), San José (2017), Aguascalientes (2018)
- Once the levelling networks are properly adjusted, a workshop about the integration of the existing height systems into the IHRS/IHRF can be planned.

# Acknowledgment



Results presented here are a joint effort of more than 50 colleagues involved in

- GGOS JWG: [Strategy for the realisation of the IHRS](#) (chair: L Sánchez)
- IAG JWG 2.2.2: [The 1 cm geoid experiment](#) (chair: YM Wang)
- IAG SC 2.2: [Methodology for geoid and physical height systems](#) (chair: J Ågren)
- ICCT JSG 0.15: [Regional geoid/quasi-geoid modelling - Theoretical framework for the sub-centimetre accuracy](#) (chair: J Huang)
- IAG JWG 2.1.1: [Establishment of a global absolute gravity reference system](#) (chair: H Wziontek)
- IAG regional sub-commissions for reference frames and geoid modelling
- IAG [Commission 2 – Gravity Field](#) (chair R Pail)
- [International Gravity Field Service – IGFS](#) (chair R Barzaghi)
- [GGOS Bureaus of Networks and Observations – GGOS-BNO](#) (chair: M Pearlman) and [Products and Standards – GGOS-BPS](#) (chair: D Angermann)